



Multivariate Statistical Analysis of Spatial and Vertical Distribution of Radiological, Mineralogical and Magnetic Parameters of Beach Sediments from North East Coast of Tamilnadu State, India

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The γ -ray spectroscopic technique has been effectively utilized to know the activity concentrations of ^{238}U , ^{232}Th and ^{40}K in different depths of north-east coast beaches of Tamilnadu state of India. The measured activity concentrations were different while compare from surface to second feet depth. FTIR exhibits the presence of eleven minerals such as quartz, feldspar in different phases, kaolinite, gibbsite, calcite, organic carbon, aragonite, palygorskite and hematite in all feet sediments. Extinction coefficient calculation shows that the amount major minerals is in the order of kaolinite < microcline feldspar < quartz. Different minerals were identified through XRD analysis. Multivariate statistical was also adopted well and the results confirm that the magnetic susceptibility play a major role in controlling the present radioactivity values. The activity concentration of ^{238}U , ^{232}Th and ^{40}K for >125 μm grain sized samples are significantly lowered when compared with the respective values of the bulk samples.

Keywords: Radioactivity, Minerals, Magnetic susceptibility, Multivariate statistical analysis, Beach sediment.

INTRODUCTION

Availability of natural radionuclides (^{238}U , ^{232}Th and ^{40}K) and their health effects to the human beings (acute leucopenia, anemia, leukemia, necrosis of the mouth, tooth fracture, pancreatic liver, hepatic, bone and kidney cancers) in various geological settings include rocks, beaches, sediments, minerals and building materials due to the local geology, geochemical behaviour and anthropogenic activities are unavoidable [1-3]. Interestingly, aforementioned presence and health effects could be related to the natural radioactivity which occupies 85% to the total global annual average ionizing radiation [4]. To assess the complete radiological risks and relevant dependence parameters of the important topographical place of communication between earthly and aquatic networks called beaches, study of natural radioactivity in concern sediments is essentially to be executed [5].

Now a day's many manmade activities and geological setting disturbance at around coastal areas especially beaches

could be possible. It is well-known that 40% of the world's population lives within 100 km of the coast. A wide majority of countries have a large percentage of their population (80-100%) living within the 100 km boundary [6] and also settlements are being concentrated within 5 km of the coastline [7]. These remarkable establishments could be available along the north-east coast of Tamilnadu state, India and were reported earlier [8]. Therefore, the determination of natural radionuclides in beaches of north-east coast of Tamilnadu is essential, which was carried out for the upper surface sediments and reported in the previous articles [9-11]. To know the complete radiological characterization of above said beaches, the present study has planned to study the radioactivity and relevant parameters in further depth (first and second feet) sediments. To achieve the goal, radioactivity data of upper surface sediment were taken for comparison [9-11].

Coastal beaches are mostly made up of sediments and sands that have been weathered and transported by water or glaciers and force of gravity [1]. Most of the coastal beach sands

or sediments contain various minerals such as quartz, feldspar, monazite, zircon and other magnetic minerals, some of them potentially contribute to the marine radioactivity [1,12]. During the mineral formations, the radionuclides could be incorporated as trace elements in their crystal lattice of minerals such as heavy, trace, magnetic and clay minerals [13]. Many studies suggest that the mineral and magnetic parameters have been recognized as simple and alternative tool for determining radioactivity level in river, marine, estuarine and fluvial sediments [14,15]. Hence, the accumulation and distribution of radionuclides depend mostly on grain size, mineral and magnetic parameters of the sediments.

The main objective of this study is to (i) determine the activity concentrations (Bq/kg) of ^{238}U , ^{232}Th , ^{40}K and calculate the absorbed dose rate (D) (nGy/h) in different depth (upper surface, first and second feet) beach sediments of north east coast of Tamil nadu state of India; (ii) identify the minerals through FTIR, XRD and measure the magnetic susceptibility in all depth sediments; (iii) find the interrelationship between radiological [^{238}U , ^{232}Th , ^{40}K and absorbed dose rate (D)], mineralogical and magnetic parameters of the beach sediment using the multivariate statistical analysis and finally, (iv) measure the activity concentrations (Bq/kg) of ^{238}U , ^{232}Th and ^{40}K in after removal of lower grain sized samples ($< 125 \mu\text{m}$) from upper surface.

EXPERIMENTAL

Study area: The study area begins from Parangipettai (Porto-novo), which is close to the Vellar river estuary and ends in Marina beach, Chennai city which is near to Coouam river estuary (lies between $11^{\circ}30'0.935 \text{ N}$ and $13^{\circ}03'0.974 \text{ N}$ latitude and $79^{\circ}46'0.279 \text{ E}$ and $80^{\circ}17'0.362 \text{ E}$ longitude). It covers more than 200 km and includes four coastal districts (Chennai, Kanchipuram, Villupuram and Cuddalore) and union territory of Pudhucherry. Other important details of the study area were clearly given in our earlier work [8].

Sample collection: From the total length (approximately 200 km), 35 places were successively chosen (S_1 to S_{35}) with distance interval of 5 to 6 km (Fig. 1). Three samples (upper surface, first and second feet) were collected manually at each sampling location during a dry period using a plastic spade (roughly 3 kg) and then stored in polyethylene bags.

Sample preparation and instruments used: The γ -ray spectrometer with NaI (Tl) crystal detector of size $3'' \times 3''$ along with 8K multichannel analyzer was used to record the γ -ray spectra. The below detectable limits (BDL) of ^{238}U and ^{232}Th radionuclide were 5.5 Bq/kg and 21.5 Bq/kg for ^{40}K , which could be determined from the background radiation spectrum for the same counting time as for sediment samples. Minerals characterizations were carried out by using the Perkin-Elmer RX1 FTIR spectrometer and Seifert (JSO-DEBYEFLEX 2002) X-ray diffractometer magnetic susceptibility measurements were carried out using a magnetic susceptibility meter, Bartington Instruments Ltd., linked to MS2B dual frequency sensor (0.47 and 4.7 KHz). Sample preparation, procedure and other technical details of instruments are already reported in our earlier works [8-10,13].

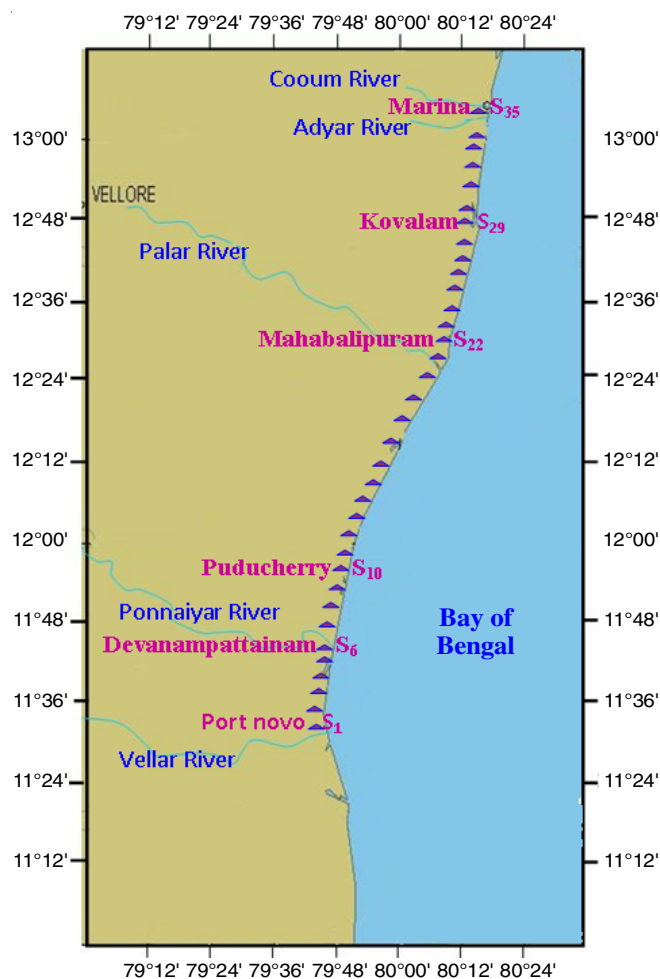


Fig. 1. Geographic location of North East coast of Tamilnadu where the sediment samples were collected

RESULTS AND DISCUSSION

Radioactivity analysis: Horizontal and vertical (upper, first and second feet samples) distribution of radionuclides

Table-1 presents the activity concentrations (Bq/kg) of ^{238}U , ^{232}Th and ^{40}K in different depth (upper, first and second feet) beach sediment samples and their distributions are shown in Fig. 2. The measured activity concentrations range from BDL to $30.42 \pm 7.90 \text{ Bq kg}^{-1}$ for ^{238}U , BDL to $218.64 \pm 8.02 \text{ Bq kg}^{-1}$ for ^{232}Th and 212.6 ± 24.68 to $423.43 \pm 26.52 \text{ Bq kg}^{-1}$ for ^{40}K in upper surface samples, BDL to $14.44 \pm 8.55 \text{ Bq kg}^{-1}$ for ^{238}U , BDL to $290.94 \pm 9.55 \text{ Bq kg}^{-1}$ for ^{232}Th and 211.03 ± 27.97 to $392.87 \pm 36.12 \text{ Bq kg}^{-1}$ for ^{40}K in first feet samples, BDL to $11.60 \pm 6.13 \text{ Bq kg}^{-1}$ for ^{238}U , BDL to $77.75 \pm 5.89 \text{ Bq kg}^{-1}$ for ^{232}Th and 206.19 ± 23.75 to $404.22 \pm 27.42 \text{ Bq kg}^{-1}$ for ^{40}K in second feet samples.

Due to mineralogy and drainage pattern of the study area, radionuclides activities are widely varied in various sites of upper, first and second feet sediments [9,10]. Activity concentration trends are not uniform from surface to deeper depth since the depositions of radioactivity minerals in the past was not uniform. However, it is very clear that the mean activity concentration of ^{238}U is decreased and ^{40}K is increased towards the deep. A random variation of ^{232}Th is observed while sea

TABLE-1
BASIC STATISTICAL DATA FOR ACTIVITY CONCENTRATIONS OF ^{238}U , ^{232}Th , ^{40}K
AND ABSORBED DOSE RATE (D) IN DIFFERENT DEPTH SEDIMENTS

	Activity concentration of radionuclides (Bq/kg)			Absorbed dose rate (D) (nGy/h)
	^{238}U	^{232}Th	^{40}K	
Upper surface				
Mean	8.31 ± 4.88	24.52 ± 4.73	275.97 ± 25.58	30.17 ± 6.11
Maximum	30.42 ± 7.90	218.64 ± 8.02	423.43 ± 26.52	156.73 ± 9.58
Minimum	BDL	BDL	212.60 ± 24.68	15.63 ± 3.50
Std. Dev	5.42	47.01	63.66	30.76
CV (%)	65.21	191.75	23.07	101.95
Skewness	3.42	3.53	0.93	3.44
Kurtosis	11.51	11.94	-0.39	11.58
First feet				
Mean	7.80 ± 4.55	38.20 ± 5.01	298.15 ± 26.57	39.11 ± 6.24
Maximum	14.44 ± 8.55	290.94 ± 9.55	392.87 ± 36.12	197.46 ± 11.41
Minimum	BDL	BDL	211.03 ± 27.97	16.16 ± 3.36
Std. Dev	2.06	53.39	51.58	33.85
CV (%)	26.47	139.75	17.30	86.54
Skewness	1.32	3.57	-0.01	3.44
Kurtosis	1.89	15.11	-1.26	14.31
Second feet				
Mean	6.64 ± 3.14	11.58 ± 3.89	299.37 ± 25.73	22.55 ± 4.88
Maximum	11.60 ± 6.13	77.75 ± 5.89	404.22 ± 27.42	64.70 ± 7.48
Minimum	BDL	BDL	206.19 ± 23.75	14.53 ± 3.37
Std. Dev	1.45	12.48	51.31	8.79
CV (%)	21.79	107.74	17.14	38.98
Skewness	1.83	4.68	0.08	3.50
Kurtosis	3.17	24.58	-0.81	15.61
Recommended level (UNSCEAR, 2000)				
	33	45	420	57

Note: Std. Dev: Standard deviation; CV: Coefficient of variability (%)

from upper to second feet depth. It is obvious from Table-1 that the high mean activity concentration of ^{238}U , ^{232}Th and ^{40}K are observed in upper surface, first and second feet samples. Uosif *et al.* [16] observed the higher concentrations of ^{40}K in deeper depths of Safaga beach, Egypt. Sutherland & de Jong [17] observed the same decreasing trend of ^{238}U towards the depth in different locations of Canada coast at the intervals of 0-15 cm and 15-30 cm, which may be due to the leachable nature of ^{238}U . The ^{40}K activity could be controlled by resistant organic matter, ordered crystalline oxide material and clay lattices [18].

In upper surface and first feet of Mahabalipuram beach (S_{22}), due to the presence of black sands enriched with mineral monazite, activity concentration of ^{238}U ($30.42 \pm 7.90 \text{ Bq kg}^{-1}$) and ^{232}Th ($290.94 \pm 9.55 \text{ Bq kg}^{-1}$) are high [16]. Likewise, as reported earlier [9-11], when compared with studied depths, in upper surface of Kovalam beach (S_{29}), activity concentration of ^{40}K ($423.43 \pm 26.52 \text{ Bq kg}^{-1}$) is high which is controlled by major minerals such as feldspar and micas [19] and higher amount of clay minerals [20].

The lowest activity concentration of ^{238}U (BDL) was found in 4 sites (S_{10} , S_{27} , S_{33} and S_{35}) of upper surface samples, 4 sites (S_3 , S_8 , S_{10} and S_{13}) of first feet samples and 9 sites (S_3 , S_4 , S_7 , S_9 , S_{10} , S_{13} , S_{16} , S_{17} and S_{33}) of second feet samples. Low concentration of ^{232}Th (BDL) was found in the upper surface of Pudhuchery beach (S_{10}), 2 sites (S_3 and S_{10}) of first feet samples and 4 sites (S_8 - S_{10} and S_{16}) of second feet samples. The lowest activity concentration of ^{40}K ($206.19 \pm 23.75 \text{ Bq kg}^{-1}$) is found

at S_{17} of upper surface. According to Jankovic *et al.* [21], the lowest activity concentrations of the radionuclides in sediments may be linked with less contribution of human activity and presence of high content of quartz.

To know the radiological hazard status basically, studied activity concentrations need to be compared with world average data (Table-1). The world averages of ^{238}U , ^{232}Th and ^{40}K in sediments were 33, 45 and 420 Bq kg^{-1} respectively [4]. From Table-1, it is confirmed that the mean activity concentrations of ^{238}U , ^{232}Th and ^{40}K in all studied depths are within the world average values. When observe the site wise, ^{238}U concentrations of all sites of all depths are lower than world average value (represented in yellow colour horizontal line) (Fig. 2). In case of ^{232}Th , three sites of upper surface, eight sites of first feet and one site from second feet exhibit the higher values than corresponding world average values. Only two sites of upper surface show the high values (^{40}K).

Univariate statistical analysis (distribution of radionuclides): The measured activity concentrations (^{238}U , ^{232}Th and ^{40}K) of samples from all studied depths have been applied into Univariate statistical analysis (Table-1) and the histograms of activity concentrations are shown in Fig. 3. The ^{238}U and ^{40}K exhibit high degree of uniformity, while ^{232}Th displays the low degree of uniformity in all studied depths. Concentrations of ^{238}U and ^{40}K radionuclides show low degree of mobility since the co-efficient of variance (CV %) values are below 100%, whereas concentration of ^{232}Th shows high degree of mobility

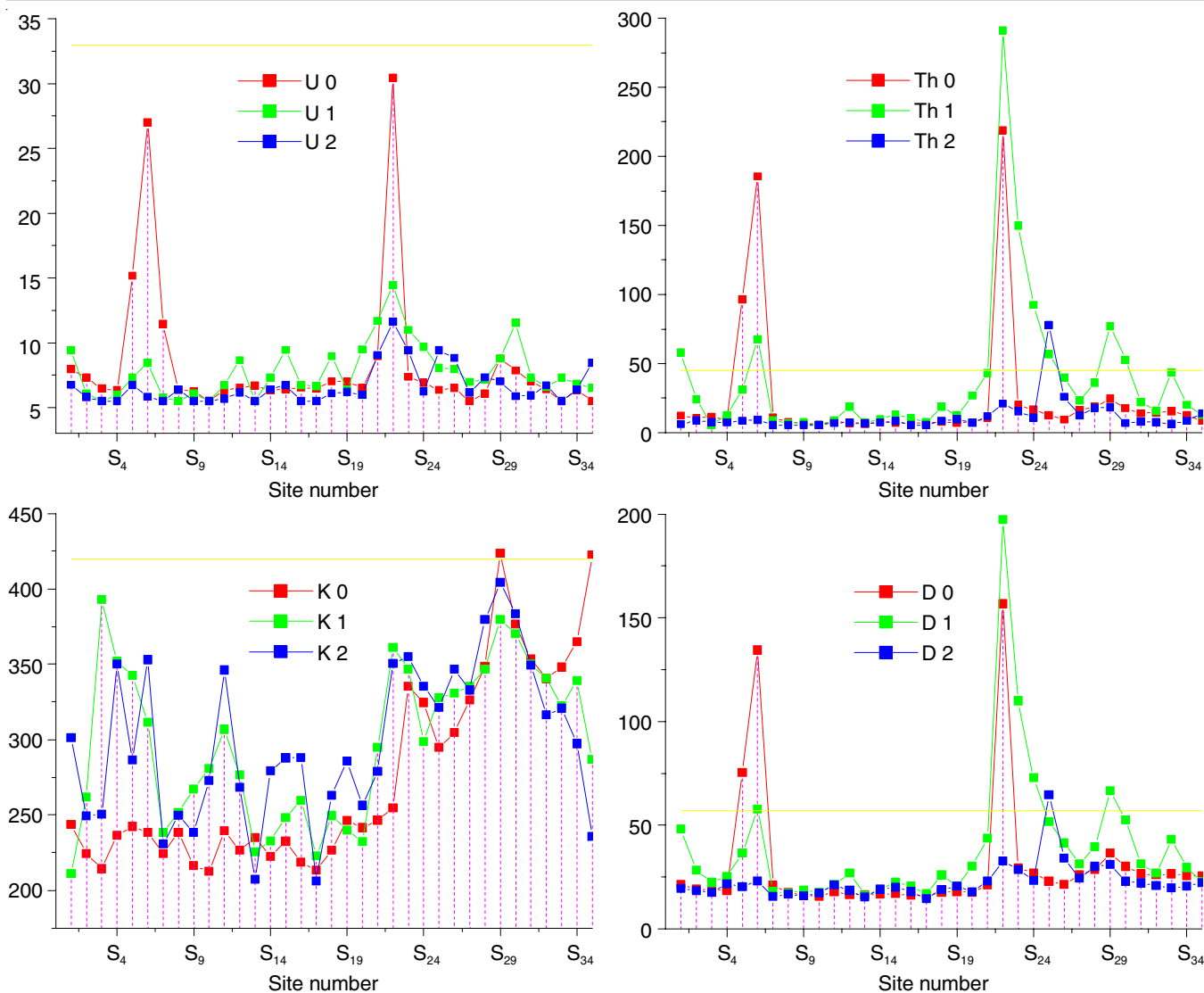


Fig. 2. Activity concentrations (Bq/kg) of ²³⁸U, ²³²Th and ⁴⁰K radionuclides in different depth beach sediment samples

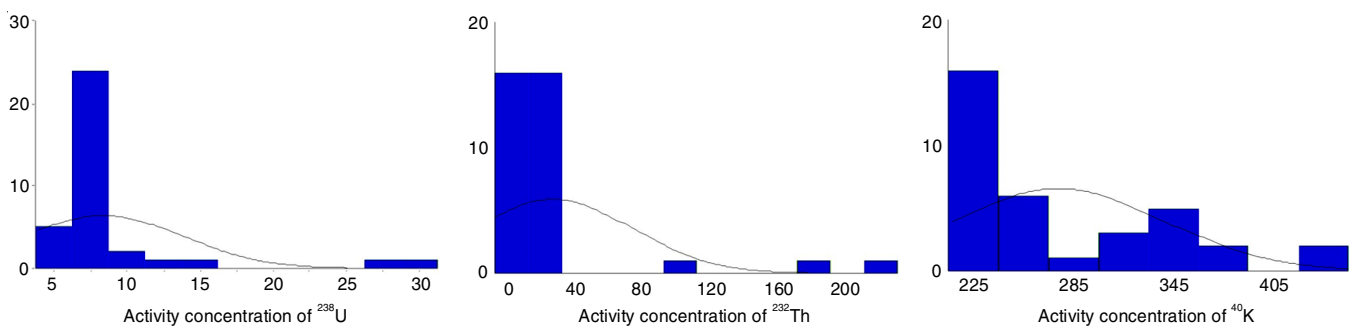


Fig. 3. Frequency distribution of ²³⁸U, ²³²Th and ⁴⁰K (upper surface samples)

since the coefficient of variance (CV %) values are above 100% [22,23] in all studied depths.

All the studied activity concentrations (except first feet of ⁴⁰K) in all feet exhibit asymmetric distributions. Their values plotted on graph do not give bell shaped form (Fig. 3; derived for upper surface) and its positive values indicate the positive skewness. Kurtosis is a measure of peakedness. Depends upon the peakedness, it is named as mesokurtic (kurtosis = zero),

leptokurtic (kurtosis = positive) and platykurtic (kurtosis = negative). In the present study, the kurtosis values of activity concentrations of ²³⁸U and ²³²Th are positive, which indicates the curve is more peaked than the normal curve (*i.e.*) leptokurtic. Negative kurtosis values are observed for ⁴⁰K concentration which indicates platykurtic (less peaked). This is due to the uneven spatial distribution of radionuclides in the sediments of north east coast of Tamilnadu [22,23].

Absorbed dose rate calculation (D): To evaluate the health risk of the present beach sediments, the total absorbed dose rate (D) (nGy/h) is calculated using the following formula [24]:

$$D = (0.462 C_U + 0.604 C_{Th} + 0.0417 C_K) \text{ nGy/h}$$

where C_U , C_{Th} and C_K are the activity concentrations (Bq/kg) of ^{238}U , ^{232}Th and ^{40}K in beach sediments, respectively. Table-1 presents the calculated values of absorbed dose rate in all foot and its spatial distribution are visualized in Fig. 2. It is ranged from 15.63 nGy h⁻¹ to 156.73 nGy h⁻¹ with an average value of 30.17 nGy h⁻¹ for upper surface samples, 16.16 nGy h⁻¹ to 197.46 nGy h⁻¹ with an average value of 39.11 nGy h⁻¹ for first feet samples and 14.53 nGy h⁻¹ to 64.70 nGy h⁻¹ with an average value of 22.55 nGy h⁻¹ for second feet samples. The mean absorbed dose rates of all depth samples are within the world average values (57 nGy h⁻¹) [4]. Three sites of upper surface and five sites of first feet samples show high absorbed dose values when compared with world average value.

Horizontal and vertical (upper, first and second feet samples) distribution of minerals

FTIR studies: The FTIR spectra were recorded for all the samples of different depth beach sediments. A representative FTIR spectrum of second feet sample of S₂₂ is shown in Fig. 4. The observed wave numbers are analyzed and the minerals are assigned (Table-2) using available literatures [9,24-35]. The minerals such as quartz, feldspar (microcline, orthoclase and albite), kaolinite, gibbsite, calcite and organic carbon are identified in upper surface samples. Almost all the minerals present in upper surface are same as in first and second feet samples with slight variations in the intensities of the peaks. Three

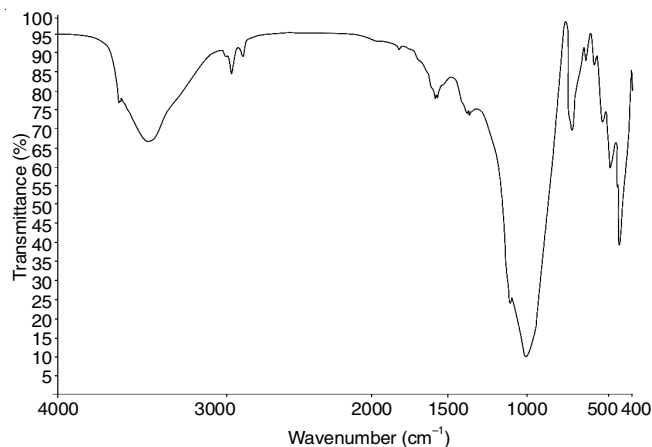


Fig. 4. FTIR spectrum of 2nd feet sample of site number 22

additional minerals (palygorskite, aragonite and hematite) were also observed in first and second feet samples.

The presence of quartz in the samples can be ascribed by the observation of the peaks in the ranges 462-458 and 514-510 cm⁻¹ due to Si-O asymmetrical bending vibrations, 694-690 cm⁻¹ due to Si-O symmetrical bending vibrations and 780-776 and 800-796 cm⁻¹ due to Si-O symmetrical stretching vibrations, while the 1084-1080 and 1164-1160 cm⁻¹ absorption region arises from the Si-O symmetrical stretching vibrations due to low Al for Si substitution. Absorption peaks in the ranges 1614-1610 and 1874-1870 cm⁻¹ also indicate the presence of quartz weathered from metamorphic origin [25,28].

The peak pertaining in the range 585-581 cm⁻¹ (O-Si-(Al)-O bending vibration), 650-646 cm⁻¹ (Al-O coordination) and 1012-1008 cm⁻¹ in the samples indicates the presence of micro-

TABLE-2
OBSERVED ABSORPTION WAVE NUMBERS AND CORRESPONDING MINERALS FROM FTIR SPECTRA FOR ALL FEET SAMPLES

Name of the minerals	Observed wavenumber (cm ⁻¹)	Site number		
		Upper surface	First feet	Second feet
Quartz	462-458	S ₁ -S ₃₅	S ₁ -S ₃₅	S ₁ -S ₃₅
	514-510	S ₁ -S ₃₅	S ₁ -S ₃₅	S ₁ -S ₃₅
	694-690	S ₁ -S ₃₅	S ₁ -S ₃₅	S ₁ -S ₃₅
	780-776	S ₁ -S ₃₅	S ₁ -S ₃₅	S ₁ -S ₃₅
	800-796	S ₁ -S ₃₅	S ₁ -S ₃₅	S ₁ -S ₃₅
	1084-1080	S ₁ -S ₃₅	S ₁ -S ₃₅	S ₁ -S ₃₅
	1164-1160	S ₁ -S ₃₅	S ₁ -S ₃₅	S ₁ -S ₃₅
	1614-1610 1874-1870	S ₁ -S ₃₅ S ₁ -S ₃₅	S ₁ -S ₃₅ S ₁ -S ₃₅	S ₁ -S ₃₅ S ₁ -S ₃₅
Microcline feldspar	585-581	S ₁ -S ₃₅	S ₁ -S ₃₅	S ₁ -S ₃₅
Orthoclase feldspar	650-646	S ₁ -S ₄ , S ₆ -S ₁₇ , S ₁₉ -S ₃₅	S ₂ -S ₄ , S ₆ -S ₁₄ , S ₁₆ , S ₁₇ , S ₁₉ -S ₂₂ , S ₂₄ -S ₃₅	S ₂ -S ₄ , S ₆ -S ₁₂ , S ₁₄ , S ₁₆ , S ₁₇ , S ₁₉ , S ₂₁ , S ₂₂ , S ₂₄ -S ₃₁ , S ₃₃ -S ₃₅
Albite feldspar	1012-1008	S ₁ , S ₁₂ , S ₁₄ -S ₁₈ , S ₂₀ -S ₃₀ , S ₃₂ -S ₃₅	S ₁ , S ₁₁ , S ₁₂ , S ₁₄ , S ₁₆ , S ₁₇ , S ₁₈ , S ₂₀ , S ₂₁ , S ₂₄ -S ₂₈ , S ₃₀ , S ₃₂ -S ₃₄	S ₁₁ , S ₁₂ , S ₁₄ , S ₁₆ , S ₁₈ , S ₂₀ , S ₂₁ , S ₂₄ -S ₂₈ , S ₃₀ , S ₃₂ -S ₃₄
Kaolinite	3400-3396	S ₁ -S ₃₅	S ₁ -S ₃₅	S ₁ -S ₃₅
	3625-3621	S ₁ -S ₃₅	S ₁ -S ₃₅	S ₁ -S ₃₅
Gibbsite	670-666	S ₁ -S ₃₅	S ₁ -S ₃₅	S ₁ -S ₃₅
Calcite	1425-1421	S ₃₂	S ₁ , S ₃ , S ₆ , S ₈ , S ₁₀ , S ₁₁ , S ₁₃ , S ₁₅ , S ₂₃ , S ₃₅	S ₆ , S ₈ , S ₁₀ , S ₁₁ , S ₁₅ , S ₂₂
Organic carbon	2924 & 2854	S ₁ -S ₃₅	S ₁ -S ₃₅	S ₁ -S ₃₅
Aragonite	1459-1455	-	S ₁ , S ₈ , S ₂₃ , S ₂₉ , S ₃₂ , S ₃₅	S ₁ , S ₈ , S ₂₉ , S ₃₂
Palygorskite	1648-1642	-	S ₁₃ , S ₁₅ , S ₂₀ , S ₃₂	S ₃ , S ₁₅ , S ₁₇ , S ₂₂ , S ₃₂
	1680-1676	-	S ₁ , S ₃ , S ₈ , S ₁₅ , S ₁₇ , S ₂₀ , S ₃₂ , S ₃₅	S ₁ , S ₃ , S ₂₉ , S ₃₂ , S ₃₅
	3615-3611	-	S ₆ , S ₈ , S ₁₀ , S ₁₃ , S ₃₂ , S ₃₅	S ₆ , S ₁₃ , S ₂₂ , S ₃₅
Hematite	533-532	-	S ₂₂ , S ₃₅	S ₂₂

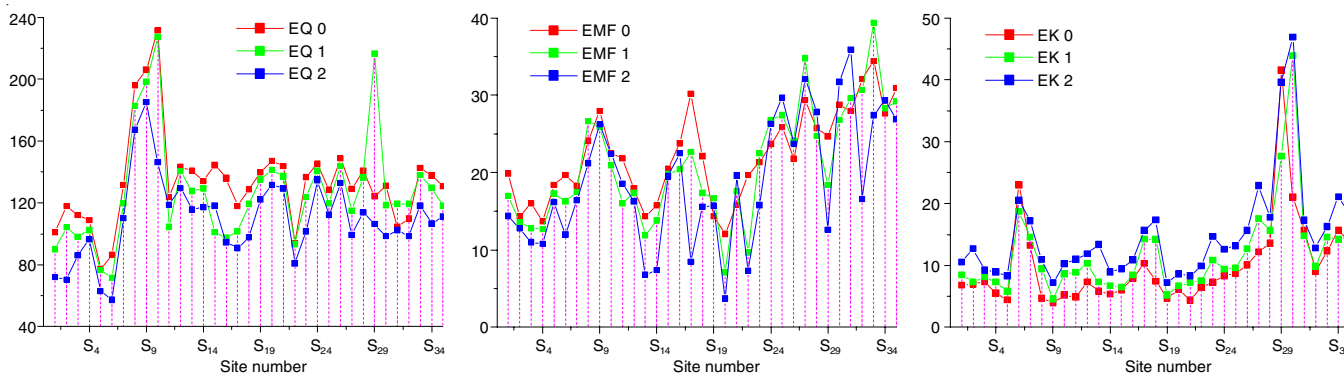


Fig. 5. Extinction co-efficient values of all depth samples

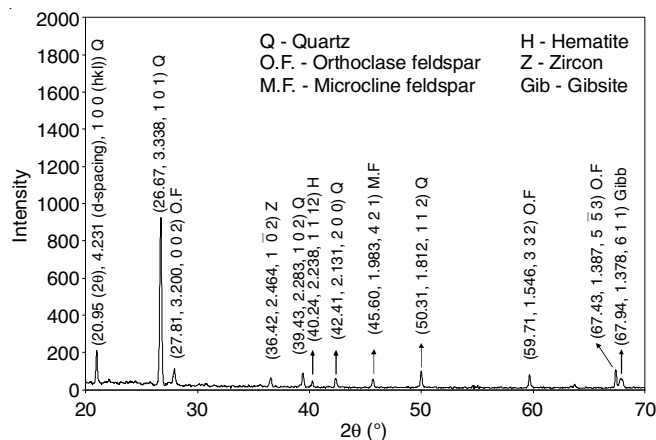
cline feldspar, orthoclase feldspar and albite feldspar ($\text{NaAlSi}_3\text{O}_8$; Na-feldspar), respectively [30]. The presence of the bands at around 3623 and 3398 cm^{-1} (OH *str.* vibration) are due to the clay mineral kaolinite. As the depth increases from upper to second feet, the intensity of the bands of kaolinite progressively increased and shifted. This suggests that the content of kaolinite increases with disorderedness as depth increases due to leaching [31].

The mid infrared region ($1500\text{--}1400\text{ cm}^{-1}$) of the spectra is dominated by the vibrational modes of carbonate ions. According to Adler & Kerr [26], when calcite mineral predominates in an aggregate the peak appears at 1423 cm^{-1} . The trihydrate aluminium mineral gibbsite is identified by the observation of the peak in the range $670\text{--}666\text{ cm}^{-1}$. Appearing peaks at 2854 and 2924 cm^{-1} in all the samples show the presence of organic carbon [30].

In present study, a peak in the range $1648\text{--}1642\text{ cm}^{-1}$ and $1680\text{--}1676\text{ cm}^{-1}$ were observed only in few sites of first feet and second feet. The presence of OH stretching band at $3615\text{--}3611\text{ cm}^{-1}$ in few sites of first and second feet samples represents the presence of palygorskite [30]. The peak observed in the range $1459\text{--}1455\text{ cm}^{-1}$ indicates the presence of aragonite mineral [10]. It is distinct from calcite (rhombohedral), which indicates its presence in some sites of first and second feet samples. Hematite is the most abundant iron mineral, which is identified in 2 sites (S_{22} and S_{35}) in first feet and only one site (S_{22}) in second feet samples, which may be due to the replacement of Fe in the samples [29].

Relative distribution of quartz, microcline feldspar and kaolinite: The extinction coefficient values of major minerals (quartz, microcline feldspar and kaolinite) were calculated [24] for all depth samples and the values are presented in Fig. 5. In all depth, it was observed that three minerals are distributed randomly. In overall view, the amount of kaolinite is lesser than feldspar and very much lesser than quartz in all the sites, which are visually indicated in Fig. 5. It is observed that the amount of quartz and feldspar is decreased and kaolinite is increased from upper to deeper depth in almost all the sites. This may be due to leaching of quartz and feldspar [31].

XRD studies: The X-ray diffractograms were recorded for all depth samples and representative diffractogram (second feet sample of S_{22}) is shown in Fig. 6. From the XRD patterns, the minerals identified in upper surface are quartz, orthoclase

Fig. 6. X-ray diffraction pattern of sample of site number S_{22} of second feet

feldspar, microcline feldspar, albite, zircon, monazite, hematite, gibbsite and goethite. In depth samples (first and second feet), quartz, orthoclase feldspar, microcline feldspar, hematite, monazite and zircon were observed. The major minerals quartz and feldspars are presented invariably in all the samples of all depth samples. The minor minerals sepiolite and the iron bearing minerals *viz.* hematite, magnetite and goethite are identified. Hence, the XRD results are agreed with the FTIR analysis.

Horizontal and vertical (upper, first and second feet samples) distribution of magnetic susceptibility: The measured magnetic susceptibility values for all depth samples are shown in Fig. 7 and are ranged from $5.5 \times 10^{-8}\text{ m}^3/\text{kg}$ to $507 \times 10^{-8}\text{ m}^3/\text{kg}$, $4.58 \times 10^{-8}\text{ m}^3/\text{kg}$ to $396.36 \times 10^{-8}\text{ m}^3/\text{kg}$ and $4.35 \times 10^{-8}\text{ m}^3/\text{kg}$ to $39.51 \times 10^{-8}\text{ m}^3/\text{kg}$ for the upper surface, first and second feet, respectively. Seven upper surface sites ($S_6, S_7, S_{21}\text{--}S_{23}, S_{25}$ and S_{33}) and only two sites (S_6 and S_{22}) from first feet samples exhibit the high magnetic susceptibility values ($>100 \times 10^{-8}\text{ m}^3/\text{kg}$) [36], which may be purely related to anthropogenic activities such as traffic effluents, emissions from vehicles, industry and fossil fuel combustions [8]. High magnetic susceptibility values are observed in all foot of Mahabalipuram (S_{22}) (Fig. 7). Spatial distributions of magnetic susceptibility in first and second feet samples almost follow the same trend observed in the upper surface. As expected, magnetic susceptibility levels fall as one move from upper to second feet. Lu & Bai [37] measured the magnetic susceptibility in surface sediments of Hangzho city, China. They also revealed the similar findings,

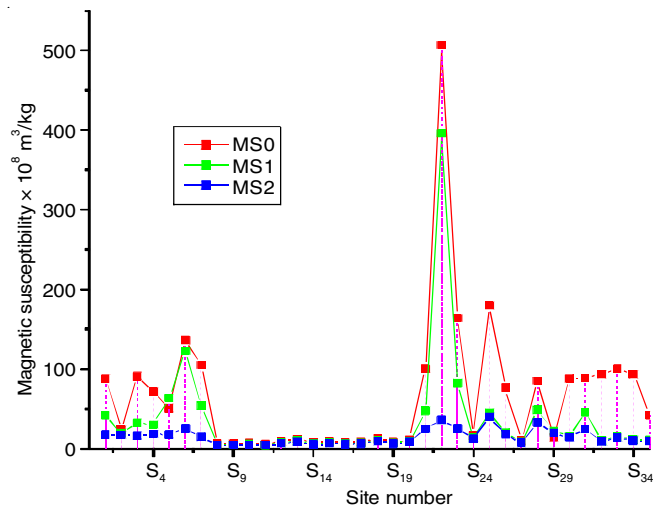


Fig. 7. Magnetic susceptibility values for all depth samples

namely that increasing levels in surface sediments are mainly related to anthropogenic activity.

Multivariate statistical analysis

Pearson’s correlation analysis: Pearson’s correlations analysis was carried out to understand the inter-relationship

between the radiological, mineralogical and magnetic parameters using the SPSS 16.0 windows software and observed results are tabulated in Table-3. A strong correlation between ²³⁸U and ²³²Th was observed in all depth samples, which may be due to the presence of monazite mineral in the study area [38]. A same strong correlation was obtained by Radhakrishna *et al.* [39], Baranwal *et al.* [40], Uosif *et al.* [16] and Ramasamy *et al.* [34] for Ullal, Chhatrapur, red sea and Kerala beach sediments, respectively. In first and second feet, a positive and poor correlation exists between ²³²Th & ⁴⁰K and ²³⁸U & ⁴⁰K, indicate that ⁴⁰K concentrations may not be related with the presence of ²³²Th & ²³⁸U bearing mineral. The same poor correlation between these radionuclides was also observed by Mohanty *et al.* [38] and Ramasamy *et al.* [34]. Existence of high positive correlation between absorbed dose rate and ²³⁸U and ²³²Th concentration in all depth samples shows that total radioactivity level is mainly controlled by these two radionuclides rather than ⁴⁰K. The same trend was observed by Ramasamy *et al.* [14] in Kerala beach sediments. Quartz and Microcline feldspar are poorly correlated, whereas Kaolinite is moderately correlated with ⁴⁰K (Table-4). Paramasivam *et al.* [13] successfully shown the dependence of radionuclide concentrations on the magnetic susceptibility in sediment. In present beach, magnetic suscepti-

TABLE-3
PEARSON’S CORRELATION ANALYSIS AMONG RADIOACTIVITY VARIABLES, EXTINCTION COEFFICIENT OF MAJOR MINERALS AND MAGNETIC SUSCEPTIBILITY IN DIFFERENT DEPTH BEACH SEDIMENTS

Upper surface								
	U0	Th0	K0	D0	EQ0	EMF0	EK0	MS0
U0	1							
Th0	0.978	1						
K0	-0.142	-0.064	1					
D0	0.971	0.997	0.015	1				
EQ0	-0.465	-0.467	-0.126	-0.477	1			
EMF0	-0.194	-0.102	0.618	-0.057	0.118	1		
EK0	0.162	0.157	0.651	0.214	-0.229	0.359	1	
MS0	0.720	0.723	0.125	0.736	-0.384	0.019	0.032	1
First feet								
	U1	Th1	K1	D1	EQ1	EMF1	EK1	MS1
U1	1							
Th1	0.781	1						
K1	0.220	0.370	1					
D1	0.786	0.998	0.422	1				
EQ1	-0.173	-0.142	-0.033	-0.142	1			
EMF1	-0.157	-0.143	0.274	-0.123	0.203	1		
EK1	0.243	0.078	0.411	0.107	0.075	0.343	1	
MS1	0.606	0.881	0.312	0.876	-0.300	-0.293	-0.052	1
Second feet								
	U2	Th2	K2	D2	EQ2	EMF2	EK2	MS2
U2	1							
Th2	0.590	1						
K2	0.311	0.266	1					
D2	0.657	0.967	0.496	1				
EQ2	-0.070	-0.012	-0.218	-0.069	1			
EMF2	0.019	0.235	0.320	0.281	0.273	1		
EK2	-0.049	0.052	0.488	0.159	-0.195	0.327	1	
MS2	0.656	0.632	0.521	0.719	-0.387	0.086	0.122	1

Note: U0, U1, U2, Th0, Th1, Th2, K0, K1, K2 respectively represents the concentrations of ²³⁸U, ²³²Th and ⁴⁰K (Bq/kg) in upper, first and second feet samples; D0, D1 and D2 respectively represents absorbed dose rate (nGy/h) in upper, first and second feet samples; EQ0, EQ1, EQ2, EMF0, EMF1, EMF2, EK0, EK1 and EK2 respectively represents relative distribution of Quartz, Microcline Feldspar and Kaolinite in upper, first and second feet samples; MS0, MS1 and MS2 respectively represents magnetic susceptibility (m³/kg) in upper, first and second feet samples.

TABLE-4
FACTOR ANALYSIS AMONG RADIOACTIVITY
VARIABLES, EXTINCTION COEFFICIENT OF MAJOR
MINERALS AND MAGNETIC SUSCEPTIBILITY IN
DIFFERENT DEPTH BEACH SEDIMENTS

Variables	Components				
	1	2	3	4	5
U0	0.798	-0.370	-0.209	0.225	0.134
U1	0.711	-0.010	0.116	0.250	-0.166
U2	0.715	0.023	0.539	-0.049	-0.054
Th0	0.820	-0.303	-0.186	0.200	0.210
Th1	0.890	-0.086	0.195	0.273	0.004
Th2	0.405	0.268	0.566	-0.496	-0.228
K0	0.273	0.838	-0.062	0.001	0.010
K1	0.547	0.505	-0.067	-0.102	0.010
K2	0.582	0.518	-0.076	-0.009	-0.169
D0	0.845	-0.237	-0.193	0.205	0.205
D1	0.902	-0.050	0.184	0.260	0.000
D2	0.543	0.358	0.507	-0.431	-0.241
EQ0	-0.558	0.160	0.545	0.526	-0.038
EQ1	-0.358	0.362	0.421	0.591	-0.294
EQ2	-0.510	0.176	0.629	0.460	-0.042
EMF0	0.009	0.755	0.008	0.067	0.541
EMF1	-0.123	0.784	0.143	-0.044	0.547
EMF2	-0.081	0.764	0.222	-0.098	0.399
EK0	0.294	0.635	-0.469	0.153	-0.322
EK1	0.223	0.711	-0.480	0.183	-0.203
EK2	0.223	0.737	-0.466	0.179	-0.279
MS0	0.892	-0.123	0.200	0.038	0.170
MS1	0.885	-0.316	0.090	0.201	0.147
MS2	0.797	0.111	0.216	-0.317	-0.147
Eigen value	8.938	5.256	2.786	1.867	1.408
Variance (%)	37.240	21.902	11.607	7.781	5.867
Cumulative (%)	37.240	59.142	70.749	78.530	84.396

Note: U0, U1, U2, Th0, Th1, Th2, K0, K1, K2 respectively represents the concentrations of ^{238}U , ^{232}Th and ^{40}K (Bq/kg) in upper, first and second feet samples; D0, D1 and D2 respectively represents absorbed dose rate (nGy/h) in upper, first and second feet samples; EQ0, EQ1, EQ2, EMF0, EMF1, EMF2, EK0, EK1 and EK2 respectively represents relative distribution of Quartz, Microcline Feldspar and Kaolinite in upper, first and second feet samples; MS0, MS1 and MS2 respectively represents magnetic susceptibility (m^3/kg) in upper, first and second feet samples.

bility plays a vital role to fix the total radioactivity level since it is well correlated with ^{238}U and ^{232}Th concentrations and absorbed dose rate in all depth samples.

Factor analysis (FA): Factor analysis (FA) is also carried out to identify a smaller number of underlying dimensions, or factors that can be used to represent relationships among a set of interrelated variables. This analysis is actually performed on the correlation matrix between different parameters (same parameters which are used in Pearson's Correlations analysis are taken) followed by varimax normalized rotation using the SPSS 16.0 windows software. Table-4 presents the factor loadings, as well as their relevant explained variance which shows that five components with eigen value > 1 , explaining 84.396 % of the total variance. The first two components show high eigen values such as 8.938 (37.24 % of total variance) and 5.256 (21.902 % of total variance). Concentrations of ^{238}U and ^{232}Th , calculated absorbed dose and measured magnetic susceptibility in all studied depths are accumulated in component 1 with high positive loadings. The ^{40}K concentration and amount

of quartz of all feet were also accumulated in component 1 with moderate positive and negative loading, respectively. It is confirm that concentrations of ^{238}U and ^{232}Th and magnetic susceptibility could be the controlling factors for total radioactivity. The parameters like ^{40}K concentration, amount of microcline feldspar and kaolinite in all feet are presented with positive loading in component 2. It shows that these parameters are least contributors for the total radioactivity. Components 3, 4 and 5 are not taken for discussions because of its lower eigen values.

From the statistical analyses, it is confirmed that magnetic susceptibility plays a major role in controlling the radioactivity values of present beach sediments. According to Paramasivam *et al.* [13] and Suresh *et al.* [41], magnetic minerals could be enriched in lower grain sized ($120\ \mu\text{m}$) samples. Moreover, several studies [20,41-45] were confirmed that the activity concentration of radionuclides decreased as the grain size of sediments decreased in different beach, coast and river sediments. Therefore, present study has intended to measure the level of radioactivity after removal of $< 125\ \mu\text{m}$ fractions from selected samples of upper surface.

Radioactivity levels of $> 125\ \mu\text{m}$ grain sized samples separated from upper surface: For this analysis, the upper surface samples were subjected into grain size separation analysis (1000 to $63\ \mu\text{m}$) by standard procedure and weight percentage of individual grain size is calculated for all the site samples (Table-5). The activity concentrations of same ^{238}U , ^{232}Th and ^{40}K (Table-6) in $> 125\ \mu\text{m}$ grain sized sediments of 15 randomly selected sampling sites from upper surface were measured by γ -ray spectroscopic analysis. It is observed that the activity concentrations are significantly lowered when compared with the respective values of the bulk samples (before the removal of $< 125\ \mu\text{m}$ grain sized particles *i.e.* upper surface) (Fig. 8) as reported in the literature.

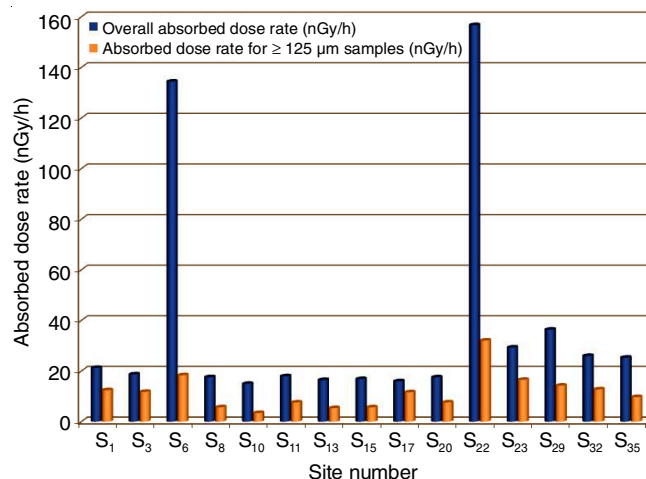


Fig. 8. Comparison of calculated absorbed dose rate for bulk and $\geq 125\ \mu\text{m}$ removed samples from selected sites from upper surface

Conclusion

The radiological, mineralogical and magnetic susceptibility measurement analyses have been successfully carried out spatially and vertically using the standard techniques. The γ -ray spectroscopic analysis shows that average of all the radiological

TABLE-5
WEIGHT PERCENTAGE OF VARIOUS GRAIN SIZED SAMPLES AND wt.% OF
MAGNETIC MINERALS OF ALL SITES OF UPPER SURFACE SAMPLES

Site number	Weight percentage of various grain sizes								Total (%)	Wt.% of $\geq 125 \mu\text{m}$
	1000 μm	710 μm	500 μm	300 μm	210 μm	125 μm	90 μm	63 μm		
S ₁	1.2	4.3	20.6	47.8	19.7	4.4	1.2	0.5	99.7	6.1
S ₂	1.7	8.6	13.8	49.4	20.1	5.0	0.7	0.3	99.6	6.0
S ₃	1.3	7.2	9.0	61.4	13.4	5.8	1.2	0.5	99.8	7.5
S ₄	1.0	2.2	10.0	52.4	26.3	6.0	1.0	0.7	99.6	7.7
S ₅	1.5	0.8	12.5	43.1	30.2	9.3	1.5	0.8	99.7	11.6
S ₆	1.0	1.4	12.8	37.0	28.2	14.9	2.6	1.8	99.7	19.3
S ₇	3.0	12.0	16.4	43.0	19.7	4.0	1.0	0.6	99.7	5.6
S ₈	8.6	18.5	26.5	27.5	14.2	3.0	1.0	0.6	99.7	4.6
S ₉	3.0	3.5	15.5	42.4	30.3	4.2	0.8	0.2	99.9	5.2
S ₁₀	3.7	6.3	15.7	44.0	24.6	3.6	1.2	0.8	99.9	5.6
S ₁₁	2.8	14.6	42.1	32.5	5.3	1.2	0.6	0.5	99.6	2.3
S ₁₂	2.0	10.3	46.4	30.7	7.2	1.5	1.0	0.5	99.6	3.0
S ₁₃	1.1	6.8	21.5	56.5	11.5	1.2	0.8	0.5	99.9	2.5
S ₁₄	1.6	8.7	30.4	44.0	12.4	1.3	1.0	0.3	99.7	2.6
S ₁₅	1.3	9.5	36.1	39.6	10.2	1.5	1.0	0.5	99.7	3.0
S ₁₆	1.2	4.8	30.5	37.0	21.6	2.4	1.5	0.5	99.5	4.4
S ₁₇	2.1	6.6	43.8	30.0	13.0	2.6	0.7	0.8	99.6	4.1
S ₁₈	1.2	5.8	39.7	29.0	19.4	2.6	1.4	0.5	99.6	4.5
S ₁₉	1.5	4.6	34.0	34.6	21.7	2.0	1.2	0.3	99.9	3.5
S ₂₀	1.0	2.4	9.4	46.7	36.2	2.7	1.0	0.4	99.9	4.3
S ₂₁	1.0	7.9	18.3	31.7	31.4	7.3	1.3	1.0	99.9	9.6
S ₂₂	1.2	10.3	12.6	17.2	34.6	21.2	1.2	1.5	99.8	23.9
S ₂₃	0.5	0.8	4.2	51.1	37.7	5.2	1.0	0.5	99.9	6.7
S ₂₄	1.3	8.2	20.6	37.6	27.4	3.0	1.3	0.4	99.8	4.7
S ₂₅	1.0	7.6	26.4	36.2	22.7	3.0	1.8	1.2	99.9	6.0
S ₂₆	1.6	10.3	27.8	35.3	19.3	3.9	1.2	0.5	99.9	5.6
S ₂₇	1.3	8.6	29.6	34.2	20.4	4.2	1.0	0.5	99.8	5.7
S ₂₈	1.4	7.2	33.1	35.6	17.2	3.2	1.0	1.0	99.7	5.2
S ₂₉	1.7	4.2	13.2	36.0	34.8	4.3	3.1	2.5	99.8	9.9
S ₃₀	1.2	3.7	16.9	32.2	36.4	4.2	2.8	1.5	98.9	8.5
S ₃₁	1.0	4.6	18.3	39.2	30.7	3.3	2.0	0.5	99.6	5.8
S ₃₂	0.7	5.7	19.3	36.0	34.1	2.5	1.0	0.5	99.8	4.0
S ₃₃	1.0	5.2	22.7	51.4	14.3	4.0	1.0	0.2	99.8	5.2
S ₃₄	0.8	4.7	26.5	44.5	18.3	3.6	1.0	0.4	99.8	5.0
S ₃₅	1.0	8.0	21.7	35.5	27.6	5.2	0.6	0.2	99.8	6.0

TABLE-6
ACTIVITY CONCENTRATIONS OF ²³⁸U, ²³²Th AND ⁴⁰K WITH
THEIR UNCERTAINTIES FOR $\geq 125 \mu\text{m}$ RANDOMLY
SELECTED SAMPLES FROM UPPER SURFACE

Site number	Activity concentration (Bq/kg)		
	²³⁸ U	²³² Th	⁴⁰ K
S ₁	5.73 ± 4.76	6.33 ± 4.39	143.74 ± 25.63
S ₃	5.68 ± 4.42	6.58 ± 4.43	124.48 ± 14.12
S ₆	8.47 ± 4.62	11.87 ± 4.94	174.36 ± 18.86
S ₈	BDL	BDL	136.78 ± 14.71
S ₁₀	BDL	BDL	82.22 ± 14.16
S ₁₁	BDL	5.76 ± 4.68	98.32 ± 16.21
S ₁₃	BDL	BDL	128.46 ± 24.38
S ₁₅	BDL	BDL	136.33 ± 24.50
S ₁₇	5.84 ± 4.46	6.72 ± 4.48	116.51 ± 24.64
S ₂₀	BDL	BDL	182.60 ± 24.68
S ₂₂	10.76 ± 4.81	28.34 ± 4.38	239.44 ± 24.72
S ₂₃	6.49 ± 4.51	6.65 ± 4.26	228.67 ± 24.70
S ₂₉	BDL	5.98 ± 4.48	254.92 ± 24.84
S ₃₂	BDL	5.76 ± 4.62	222.62 ± 25.64
S ₃₅	BDL	BDL	233.49 ± 25.14

parameters (concentrations of ²³⁸U, ²³²Th and ⁴⁰K and calculated absorbed dose rate) were within the recommended level in all studied feet. However, some extreme values were observed in upper surface and first feet of S₆ and S₂₂ due to the presence of black sands. Various activity concentration trends are observed from surface to second feet depth. However, mean activity concentration of ²³⁸U was decreased, while ⁴⁰K was increased towards the deep with random variation of ²³²Th. The FTIR and XRD analyses exhibit the presence of eleven minerals in all feet sediments. The amount of quartz and feldspar is decreased and kaolinite is increased from upper to deeper depth (first and second feet) due to weathering and leaching of quartz and feldspar. Due to anthropogenic activities, upper surface samples have higher magnetic susceptibility value when compared with depth samples. Statistical analysis confirmed that magnetic susceptibility plays a major role in controlling the radioactivity values of present beach sediments. The activity concentration of ²³⁸U, ²³²Th and ⁴⁰K of $> 125 \mu\text{m}$ grain sized

sediments from upper surface are significantly lowered when compared with the respective values of the bulk samples.

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CONFLICT OF INTEREST

The authors declare that there is no conflict of interests regarding the publication of this article.

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