

Chemical Characteristics and Formation of Karst Water in Guilin Area, China

Hui Qian^{1,2,*}, Jianhua Wu^{1,2} and Peiyue Li^{1,2}

¹School of Environmental Science and Engineering, Chang'an University, No. 126 Yanta Road, Xi'an 710054, Shaanxi, P.R. China ²Key Laboratory of Subsurface Hydrology and Ecology in Arid Areas, Ministry of Education, No. 126 Yanta Road, Xi'an 710054, Shaanxi, P.R. China

*Corresponding author: Fax: +86 29 82339952; Tel: +86 29 82339327; E-mail: qianhui@chd.edu.cn

(*Received*: 28 November 2011; *Accepted*: 19 September 2012) AJC-12156

In present study, the chemical characteristics and evolution law of the Guilin karst water are reported. The chemical compositions of 74 karst water samples and one local rain water sample were compared and analyzed. The dissolved amounts of minerals and gas in the 74 karst water samples were calculated using mass balance approach. The forming environments of the karst water were analyzed and discussed. The results show that the karst water in this region is formed by the dissolutions of gypsum, calcite, dolomite and atmospheric CO_2 on the basis of the chemical compositions of the local rain water and the Guilin karst water is formed in a relatively open environment.

Key Words: Karst water, Dissolution of mineral and gas, Mass balance calculation, Forming environment, Water chemistry.

INTRODUCTION

Karst regions are characterized by limestones and other soluble rocks at or near land surface that have been modified by solution erosion¹. Interest in karst hydrology has increased greatly in the past decades, which has resulted in international exchange of many topical karst problems. The history of the research on karst water in China can be traced back to the 1960s, when a national karst conference was held by the Geological Society of China in Guilin². Since then a number of scholars have focused on karst water research at different levels in China. Gu³ put forward several experiments about shallow karst water service to agricultural production. At this time, some international scholars have also noticed the interesting research. Pitty⁴ proposed an approach of correlating both solute fluctuations and flow rate of karst water with climatic data to study karst water. In their work, nearly 100 karst water samples were collected and analyzed for determining the quantity of dissolved calcium carbonate in the limestone areas of the southern and central Pennines. In the 1970s, some specific investigations and calculations on karst water were carried out in some river basins⁵⁻⁷ and the runoff features of karst water and the effects of karst features on water circulation were studied^{8,9}. Entering into the 1980s, more studies on karst problems were conducted. These studies nearly covered every aspect of karst study such as chemical characteristics of karst water¹⁰⁻¹², chemical classification of karst water¹³, karst water pollution¹⁴, movement and dynamics of karst water¹⁵, karst

landscape 16,17 , the formation of karst water $^{18-20}$ and karst water resource evaluation and management²¹⁻²⁴. Since the 1990s, the studies on various aspects of karst related issues have become more complex and with the popularity of computers, numerical modeling on kasrt water movement has become more and more $common^{25-28}$.

At present, there are typically two ways for karst phenomenon research: one is the direct way which studies directly the changing characteristics of dissolvable minerals when water flows through them and the other is the indirect way which discusses the changing characteristics of dissolvable minerals indirectly by studying the physicochemical property changes of water which flows through these minerals. Actually, the latter way is to calculate the dissolution and precipitation of minerals by discussing the physicochemical property changes of the water flowing through the dissolvable rocks. This paper attempts to discuss the dissolution and precipitation of minerals in Guilin karstic region by the latter way.

EXPERIMENTAL

Study area: Guilin area, one of the most popular areas for karst research in China and the world, is situated between longitude 110º09' -110º42'E and latitude 24º40' -24º40' N, and covers an area of 7420 km² . It is covered by limestone all over the area and is the most typical and developed area for karst landforms and karst groundwater in China. The first international karst research institute is located in the area. The area is dominated by subtropical monsoon climate and is characterized 1542 Qian *et al. Asian J. Chem.*

ŕ

Vol. 25, No. 3 (2013) Chemical Characteristics and Formation of Karst Water in Guilin Area, China 1543

63	7.90	17.00	2.84	9.61	205.06	0.40	0.30	68.54	0.73	0.00
64	8.44	12.50	2.13	11.53	180.65	0.40	0.00	67.74	0.00	7.20
65	8.37	13.00	3.55	12.49	124.53	0.30	0.00	45.69	0.49	2.40
66	7.41	18.00	2.13	7.68	12.42	0.20	0.00	103.81	3.19	0.00
67	7.40	18.00	2.13	7.68	280.68	0.20	0.00	91.78	1.94	0.00
68	7.48	18.20	2.13	7.68	207.46	0.20	0.00	70.94	0.97	0.00
69	7.72	16.50	2.84	6.72	170.86	0.62	0.29	53.71	2.19	0.00
70	7.74	19.50	2.84	9.61	163.53	50.50	0.00	55.31	1.22	0.00
71	7.66	18.20	2.13	24.02	180.62	0.40	0.00	67.33	0.49	0.00
72	7.46	19.20	3.55	23.06	263.61	0.65	0.00	92.99	2.92	0.00
73	7.62	15.50	1.42	62.44	153.76	0.40	0.00	73.75	1.70	0.00
74	7.50	17.50	1.42	24.98	180.61	0.30	0.00	63.33	2.67	0.00
Mean	7.71	17.89	3.66	10.28	191.51	0.67	0.23	65.27	2.05	1.23
Min	7.16	7.3	0.00	0.00	12.42	0.00	0.00	18.04	0.00	0.00
Max	8.76	31	21.98	62.44	336.89	10.83	2.55	109.02	8.03	14.4
SD	0.38	3.46	3.16	9.12	64.18	5.96	0.36	20.98	1.70	2.95

Mean, Min, Max, and SD are mean, minimum, maximum and standard deviation of indices of karst water, respectively.

by hot and rainy summer, cold and dry winter. The average annual temperature is 18-19 ºC and increases from north to south. The average annual precipitation in the area is 1897.3 mm with over 60 % concentrated during April to July²⁹. The main minerals forming the rocks are dolomite, calcite and gypsum which can easily dissolve under specific conditions. The world famous scenery in the area is formed by the dissolution of these minerals. Rain water, the most important source for karst water, is essential for the forming of karst landforms. All the karst water bodies including karst surface water and karst groundwater are formed on the basis of rain water through the water-rock interaction.

RESULTS AND DISCUSSION

Chemical characteristics of Guilin karst water: To carry out the study, a total of 74 karst water samples and a rain water sample were used. Because of the inadequate of available data, all these data used in the paper were abstracted from the monograph³⁰. The chemical analysis results and the statistical analyses of karst water samples were included in the Table-1 in which the rain water chemical analysis results were also listed. The temperature of rain water is the local average air temperature. Compared the chemical compositions of local rain water with those of karst water, it can be found that the karst water possesses the following characteristics:

(i) Compared with the rain water, the concentrations of Cl⁻, Na⁺ and K⁺ in karst water do not change significantly. The concentrations of Cl^- , Na⁺ and K⁺ in rain water are 3.50, 0.20 and 0.00 mg/L and the mean values of the three chemical parameters in karst water are 3.65, 1.35 and 0.23 mg/L. The highest concentration of Na⁺ is 10.83 mg/L observed in sample 47 and the lowest is 0. The concentrations of K^+ are lower than 1.0 mg/L in all karst water samples except in sample 22 which possesses the highest K^+ concentration of 2.55 mg/L.

(ii) The increases of the concentrations of SO_4^2 , $HCO_3^$ in karst water samples are significant compared with those in the rain water sample. In the rain water, no SO_4^2 was detected, but in karst water samples the mean concentration of SO_4^2 reaches to 10.28 mg/L and its biggest concentration is 62.44 mg/L. The maximum concentration of $HCO₃⁻$ in karst water samples is 336.89 mg/L and the mean concentration is 187.46 mg/L which is over 31 times higher than that in the rain water. A significant increase of Ca^{2+} content in karst water was also observed. The mean Ca^{2+} content in karst water is 40.79 times of that in rain water.

(iii) A slight increase of Mg^{2+} concentration in karst water was observed. The mean $\overline{M}g^{2+}$ concentration in karst water is 2.05 mg/L which is 51.26 times of that in the rain water. The concentration of $CO₃²$ in karst water is mainly controlled by pH. Generally speaking, only when pH value is higher than 8.0 can the $CO₃²$ exist and be detected in water.

Based on the geology, hydrogeology of the study area and the above comparison, it can be concluded that the changes of the chemical compositions in karst water relative to those in rain water are resulted from the dissolution of calcite, dolomite, gypsum and carbon dioxide $(CO₂)$.

Dissolution calculation of the minerals and gases in karst water: Since the chemical changes of the karst water relative to the local rain water are resulted from the disolution of calcite, dolomite, gypsum and $CO₂$, it is of great concern to determine how much dissolution of these minerals is needed to form current chemical constituents of karst water on the basis of the rain water chemistry. According to the elements mass balance among minerals, gases and water, the amounts of minerals and gases dissolved in karst water can be calculated.

There are four main elements in gypsum, calcite, dolomite and $CO₂$ that are our most concern, namely, S, C, Ca and Mg. Assume that X_1 mmol/L gypsum, X_2 mmol/L CO₂, X_3 mmol/ L calcite and X_4 mmol/L dolomite are dissolved to form the existing chemical compositions of the karst water, then the dissolved amounts for element S, C, Ca and Mg are X_1 , $(X_2 +$ $X_3 + 2X_4$, $(X_1 + X_3 + X_4)$ and X_4 , respectively. At the same time, the above dissolution makes the element S, C, Ca and Mg contents in water solution also increase. If dS, dC, dCa and dMg are used to denote the increase of element S, C, Ca and Mg per liter water, respectively, the following relationships according to the principle of mass balance should be met:

$$
X_1 = dS \tag{1}
$$

$$
X_2 + X_3 + 2X_4 = dC \tag{2}
$$

$$
X_1 + X_3 + X_4 = dCa
$$
 (3)

$$
X_4 = dMg \tag{4}
$$

1544 Qian *et al. Asian J. Chem.*

Ξ

Vol. 25, No. 3 (2013) Chemical Characteristics and Formation of Karst Water in Guilin Area, China 1545

64	0.120	16.33	1.451	63.86	1.535	153.51	-0.002	-0.31	1.9475
65	0.130	17.69	0.992	43.63	0.953	95.34	0.019	3.46	2.0008
66	0.080	10.88	2.417	106.34	2.344	234.40	0.131	12.14	1.9272
67	0.080	10.88	2.249	98.97	2.095	209.53	0.079	14.57	1.9981
68	0.080	10.88	1.610	70.86	1.615	161.48	0.039	7.14	1.9516
69	0.070	9.52	1.380	60.73	1.143	114.32	0.090	16.49	2.0438
70	0.100	13.61	1.291	56.79	1.194	119.35	0.049	9.05	1.9991
71	0.250	34.03	1.451	63.84	1.374	137.43	0.019	3.44	2.0277
72	0.240	32.67	2.059	90.58	1.924	192.45	0.120	22.08	1.9510
73	0.650	88.45	1.200	52.79	1.084	108.42	0.069	12.73	1.9814
74	0.260	35.39	1.470	64.67	1.173	117.35	0.110	20.17	2.0554

In the above equations, dS, dC, dCa and dMg can be obtained according to water chemistry analysis results. Solving the equations of l, 2, 3 and 4, the amounts of dissolved gypsum, calcite, dolomite and $CO₂$ can be obtained. This method is called mass balance calculation (MBC). The calculated results of MBC for the 74 karst water samples are shown in Table-2. Ratio in Table-2 denotes the ratio of the dissolved element C to the sum of dissolved Ca and Mg by calcite and dolomite which in value equals to $(X_2 + X_3 + 2X_4)/(X_3 + 2X_4)$.

Table-2 reveals that all the samples are formed by gypsum, calcite, dolomite and $CO₂$ on the basis of rain water except sample 64 which is formed by the dissolution of gypsum, calcite and $CO₂$, and the precipitation of dolomite (denoted by minus number for dolomite, -0.31 mg/L). The average amounts of dissolved gypsum, calcite, dolomite and $CO₂$ are 14.56, 140.08, 15.41 and 65.71 mg/L. The largest amount for dissolved gypsum is 88.45 mg/L observed in sample 73. The largest amount for dissolved $CO₂$ is observed in sample 62, being 121.40 mg/L. The biggest amounts for dissolved calcite and dolomite are, respectively observed in samples 9 and 32 and they are 243.49 and 61.25 mg/L, respectively.

If the amounts of dissolved $CO₂$ in the processes of forming the chemical compositions of the local surface water and groundwater in Guilin area are estimated, some very meaningful results can be obtained. The Guilin area covers an area of 7420 km² and the annual average rainfall in the area is 1897.3 mm. Seventy percent of the rainfall forms surface water and groundwater, and the total surface water and groundwater formed by the precipitation is 9.8546×10^9 m³ per year. If the mean amount of dissolved $CO₂$ is taken into account, then 6.475×10^5 t CO₂ will dissolve into the rain water to form the local surface water and groundwater each year. Even if the minimum amount of dissolved $CO₂$ of the 74 karst water samples is used to calculate the annual dissolved $CO₂$ amount, the result is 1.614×10^5 t per year. This result is of particular importance for protecting earth's environment, especially in these years when the emission of $CO₂$ is increasing.

The dolomite precipitation occurred in the sample 64, seems impossible according to its chemical compositions. It may because that there is no Mg^{2+} in sample 64 and when the average Mg^{2+} concentration of rain water which is higher than that in sample 64 is used for calculation, incorrect result may arise.

Forming environment of karst water in Guilin: There are two typical environments where dolomite and calcite dissolve into the rain water, namely, closed environment (closed system) and open environment (open system). In a closed system, calcite and dolomite usually dissolve into water by the following formulas:

$$
CaCO3 \longrightarrow Ca2+ + CO32
$$

\n
$$
CaMg(CO3)2 \longrightarrow Ca2+ + Mg2+ + 2CO32
$$

In closed systems, there is no $CO₂$ recharge during the reactions and the ratio of the dissolved element amount of C to the sum of Ca and Mg calculated by calcite and dolomite dissolution (ratio in Table-2) should be 1. Whereas, in an open system there is continuous $CO₂$ recharge during the reactions and the reactions are expressed as follows:

 $CaCO₃ + CO₂ + H₂O \longrightarrow Ca²⁺ + 2HCO₃$ $CaMg(CO₃)₂ + 2CO₂ + 2H₂O \implies Ca²⁺ + Mg²⁺ + 4CO₃$ The ratio in the open system should be equal to 2 according

to the above reaction formulas. The last column of Table-2 indicates that the mean value of ratio for the 74 karst water samples in Guilin is 1.9469 and

the minimum value is 1.6046. So it can be conclude that the karst water in Guilin is formed in an open environment.

Conclusion

The chemical compositions of the karst water in Guilin are formed by dissolution of calcite, dolomite, gypsum and $CO₂$ on the basis of the chemical compositions of the local rain water. Mass balance calculation can be used to calculate the dissolved amounts of the minerals and gases of the karst water relative to the local rain water in Guilin. The least amount of dissolved CO₂ estimated is 1.614×10^5 t per year and average is 6.475×10^5 t per year, which is particularly important and useful for $CO₂$ storage and improving and protecting the earth's environment, and is also of significance in controlling $CO₂$ emission and global climate warming. Overall, the karst water in Guilin is formed in an open environment.

ACKNOWLEDGEMENTS

The research was supported by Special Fund for Basic Scientific Research of Central Colleges, Chang'an University (CHD2011TD003), the Doctor Postgraduate Technical Project of Chang'an University (CHD2011ZY025 and CHD2011ZY022) and the National Natural Science Foundation of China (41172212).

REFERENCES

- 1. H.E. Legrand and V.T. Stringfield, *J. Hydrol*., **20**, 97 (1973).
- 2. Y.R. Lu, *Geological Rev*., **12**, 13 (1966) (in Chinese).
- 3. J.G. Gu, *Water Resour. Hydropower Eng*., 51 (1965) in Chinese.
- 4. A.F. Pitty, An Approach to the Study of Karst Water, Illustrated by Results from Poole's Cavern, Buxton, Occasional Papers in Geography No. 5, University of Hull (1966).
- 5. S.N. Davis, *Earth-Sci. Rev*., **8**, 324 (1972).
- 6. Survey Institute of North China, Geotechnical Investigation and Surveying, No. 12, pp. 45-56 (1975) (in Chinese).
- 7. H.C. Liu, *Hydrogeol. Eng. Geol*., 60 (1979) (in Chinese).
- 8. V.T. Stringfield and H.E. LeGrand, *J. Hydrol*., **14**, 139 (1971).
- 9. V.T. Stringfield, J.R. Rapp and R.B. anders, *J. Hydrol*., **43**, 313 (1979).
- 10. H.X. Peng and Y.K. Wu, *Carsologica Sin*., **2**, 117 (1983) (in Chinese).
- 11. Y.X. Zhou, *Hydrogeol. Eng. Geol*., 6 (1987) (in Chinese).
- 12. J. Boulègue, M. Benedetti and P. Bildgen, *Appl. Geochem*., **4**, 37 (1989).
- 13. K.S. Wang, *Chongqing Environ. Protection*, 28 (1987) (in Chinese).
- 14. N. Simmleit and R. Herrmann, *Water Air Soil Pollut*., **34**, 79 (1987).
- 15. E.B. Cao, *Carsologica Sin*., **5**, 15 (1986) (in Chinese).
- 16. M.D. Yang, *Carsologica Sin*., **1**, 81 (1982) (in Chinese).
- 17. R.A.L. Osborne and D.F. Branagan, *Earth-Sci. Rev*., **25**, 467 (1988).
- 18. L.H. Song, Y.G. Zhang, J.F. Fang and Z.X. Gu, *J. Hydrol*., **61**, 3 (1983).
- 19. D. Buhmann and W. Dreybrodt, *Chem. Geol*., **48**, 189 (1985).
- 20. D. Buhmann and W. Dreybrodt, *Chem. Geol*., **53**, 109 (1985).
- 21. C.Y. Ha, H.Y. Wang and P. Yang, *Geotechnical Investigation Surveying*, 66 (1980) (in Chinese).
- 22. Z.S. Liao, *Carsologica Sin*., **4**, 101 (1985) (in Chinese).
- 23. J.B. Wang, *J. Taiyuan Univ. Technol*., 69 (1985) (in Chinese).
- 24. W.P. Wang, *Jiangsu Geol*., 33 (1986) (in Chinese).
- 25. W. Dreybrodt, *J. Geol*., **98**, 639 (1990).
- 26. Z.S. Li, *Coal Geol. China*, **7**, 55 (1997) (in Chinese).
- 27. E.S. Lee and N.C. Krothe, *Chem. Geol*., **179**, 129 (2001).
- 28. I.L. Dryden, L. Márkus, C.C. Taylor and J. Kovács, *J. Royal Statistical Soc.: Series C (Appl. Statistics)*, **54**, 673 (2005).
- 29. Z.Z. Gong, *Geological Rev*., **33**, 347 (1987).
- 30. X.W. Zhu, X.Y. W, D.H. Zhu, Z.Z. Gong and H.R. Qin, Guilin Karst Landforms and Caves Research, Beijing: Geological Publishing House (1988) (in Chinese).